



# MARINE METALLIFEROUS SEDIMENTS: A POTENTIAL FUTURE SOURCE OF ESSENTIAL METALS?

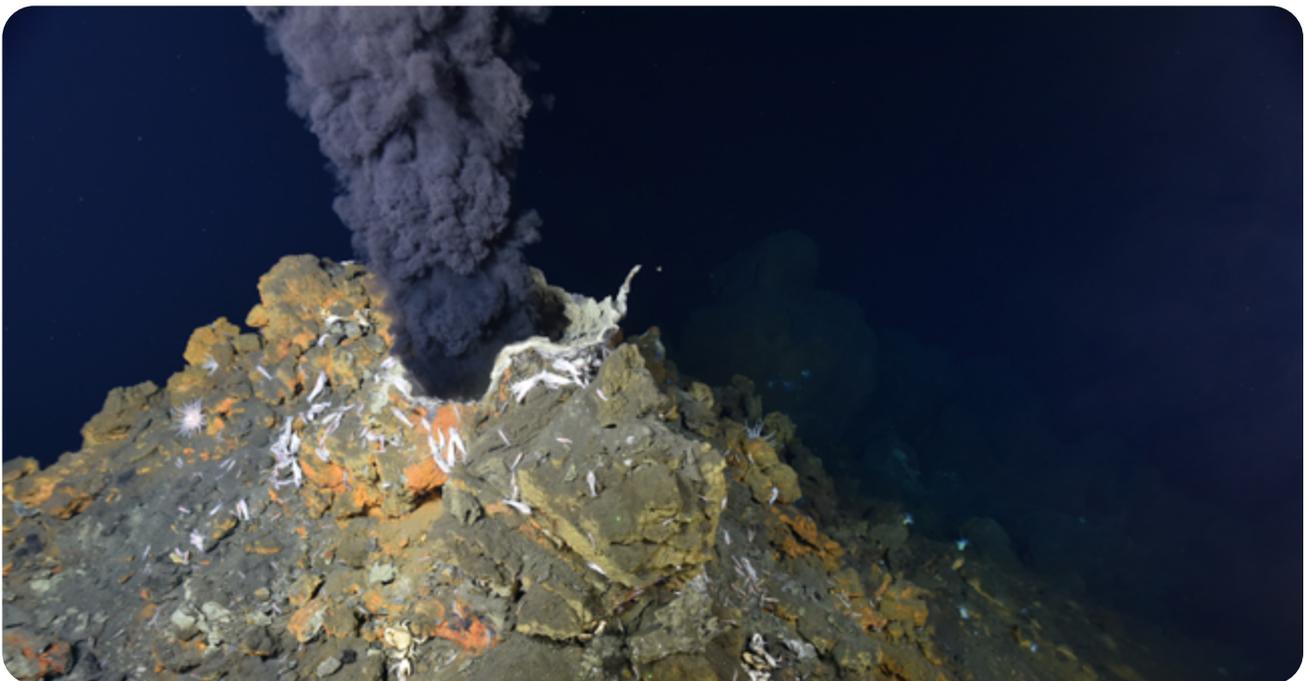
## INTRODUCTION

Marine metalliferous sediments are widespread geological formations on the ocean floor that may contain significant concentrations of potentially economically valuable metals. These sediments may be enriched in various metals and elemental groups, depending on a combination of sedimentation and seabed chemistry over time, including rare-earth elements (REE).

Geological processes, such as regional hydrothermal activity, sedimentary processes and the slow accumulation of dissolved metals in seawater, influence the formation of marine metalliferous sediments. These deposits are particularly interesting due to their potential as a future source of essential metals for various industries, including electronics and renewable energy technologies.

Researchers and governments are exploring the feasibility of extracting and utilizing marine metalliferous sediments. Their studies contribute to our understanding of Earth's geological processes over time and the processes in the seabed environment.

Where they occur in the Area (defined as the seabed and subsoil beyond the limits of national jurisdiction), marine metalliferous sediments are one of its resources subject to regulation by the International Seabed Authority (ISA) under the United Nations Convention on the Law of the Sea (UNCLOS) and its 1994 Implementing Agreement. However, they are not addressed in the Exploration regulations or the Exploitation regulations currently under negotiation.

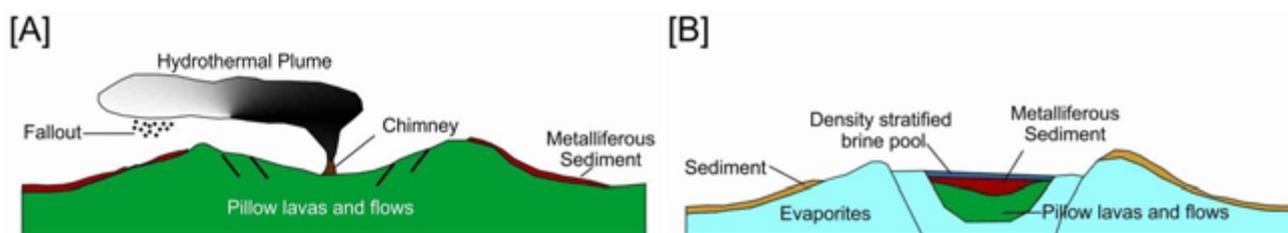


## HISTORICAL CONTEXT, GEOLOGICAL DEFINITION AND FORMATION

Marine metalliferous sediments were first identified during the HMS Challenger expedition between 1873 and 1876 (Murray and Renard, 1891). They represent potential metal resources and a tool for exploring hydrothermal vent sites (Gurvich, 2006) and vent districts. Unconsolidated, metal-enriched sedimentary deposits are the modern analogues of lithified and unlithified ancient marine sediment deposits of the Paleo-oceans, often found as extensive exhalite horizons in ophiolite belts such as Oman and Cyprus (Fleet and Robertson, 1980; Robertson and Hudson, 1973) and in older crustal rifting regions (Hoggard et al., 2020). Metalliferous sediments form through plume fall out and mass wasting related to hydrothermal vent activity (James and Elderfield, 1996; Mills and Elderfield, 1995) and/or uptake of dissolved metals from seawater (Boström and Peterson, 1969; Gurvich, 2006) or are related to isolated hot brine pools (Degens and Ross, 1969). Because their formation may be directly linked to the degradation of polymetallic sulphide (PMS) deposits, the metal content of metalliferous sediments is also strongly dependent on the fluid composition at each hydrothermal vent site. In general, metalliferous sediments are enriched in iron and manganese that form oxides and (oxy-) hydroxides, trace amounts of Fe-Cu-Zn-Pb-sulphide minerals (such as pyrite, marcasite, galena, sphalerite, chalcopyrite), silica polymorphs, including quartz and amorphous silica, barite and many more. Transitional metals, such as cobalt, nickel, gold, silver, vanadium, REEs and uranium, can be enriched locally (German et al., 1993; Takaya et al., 2018).

### FORMATION

Deep-sea metalliferous sediments generally combine varied modes of formation with components and inputs from multiple sources. They often have multiple formation processes and vary in composition. Most reported occurrences and accumulations are associated with hydrothermal activity along oceanic spreading centres, where hydrothermal fluids discharge (Figure 1, frame A). Once the high-temperature hydrothermal fluids vent and mix with cold near-bottom seawater, very fine-grained particles precipitate as plume fall out (Dymond, 1981). In addition, due to weathering and mass-wasting processes on the seafloor, massive sulphide mounds may form from clastic sulphide debris (Metz et al., 1988; Mills and Elderfield, 1995; Halbach et al., 1998).

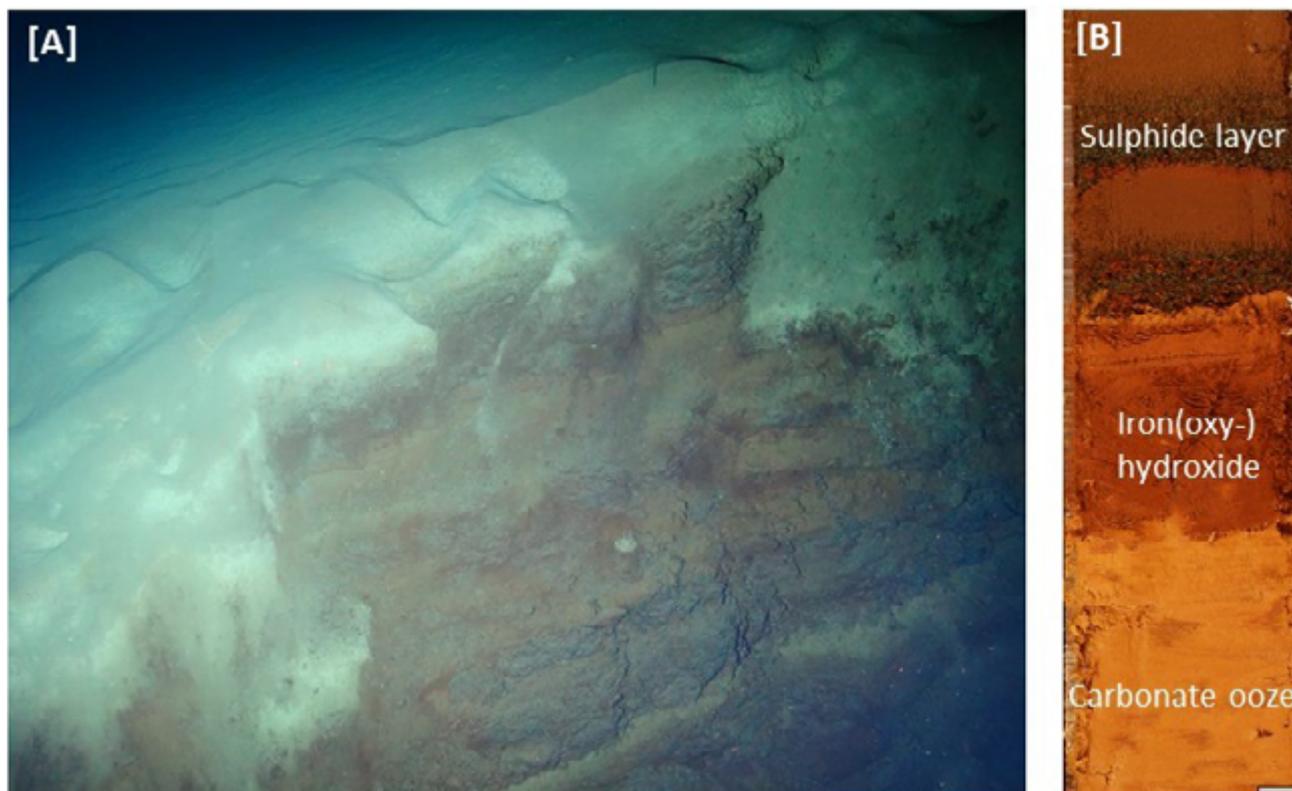


**Figure 1.** Schematic diagram showing the difference between the formation of metalliferous sediments in mid-ocean ridge settings (A) and the Red Sea (B). Most of these sediments accumulate in small depressions, so-called sediment ponds (Rona et al., 1993), where they form alternating layers of red-brown mud with intercalated layers of carbonate, silicic ooze, or sulphide debris (Figure 2). Low sedimentation rates facilitate metal scavenging from seawater. Metals are either incorporated into or adsorbed onto particles predominantly composed of iron and manganese (oxy-) hydroxides (Boström and Petersen, 1969; Marchig and Erzinger, 1986; Gurvich, 2006; Li and Schoonmaker, 2014), similar to the formation processes of polymetallic manganese nodules (PMN).

**Source:** Modified from Zierenberg (1990)

In contrast, metalliferous sediments of the Red Sea form from different mechanisms (Figure 1, frame B). Rifted graben structures with submarine brine pools originating from the exhalation of high-salinity brines occur along the central axes of the Red Sea. Sulphides, such as pyrite, sphalerite, galena and chalcopyrite, precipitate from a lower brine layer, whereas iron-(oxy-)-hydroxides and amorphous silica

precipitate from an upper brine layer, with these minerals later forming clay mineral phases by diagenetic processes (Weber-Diefenbach, 1977). The brines' high salinity, elevated temperature and heavy metal concentrations ensure a hostile environment for organisms (Watson and Waterbury, 1969). Thus, bioturbation, sulphate reduction and oxidation processes are inhibited, resulting in undisturbed accumulations that are further protected from weathering by overlying anoxic brines. The overall metalliferous sulphide accumulations cover an area of 60 km<sup>2</sup> and can reach a thickness of up to 20 m (Zierenberg, 1990).

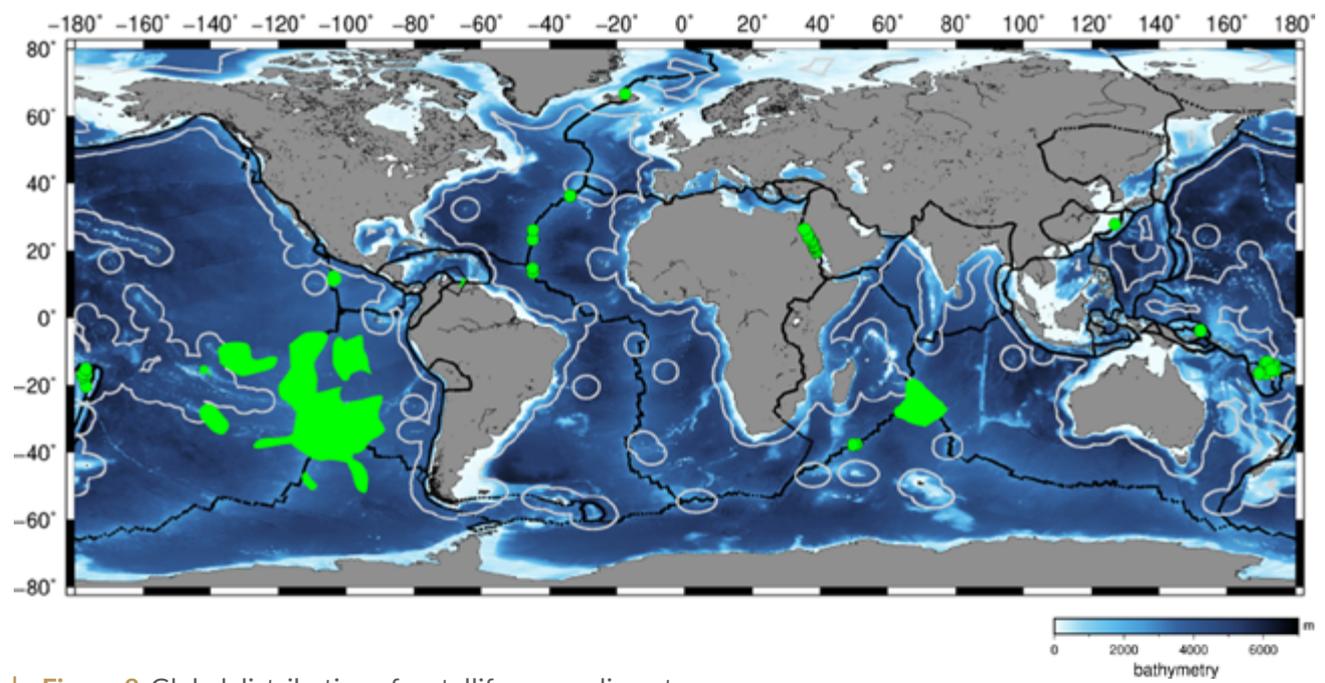


**Figure 2.** Sediment accumulation from the Transatlantic Geotraverse TAG hydrothermal vent area

**Legend:** An outcrop of metalliferous sediments at a fault scarp that is overlain by carbonate ooze (A), sulphide layers in iron-rich hydrothermal sediment that alternate with carbonate ooze (B) **Image credit:** National Oceanography Centre; ISA, 2021

## OCCURRENCES

Significant accumulations of metalliferous sediments are found in mid-ocean ridge settings, where their precipitation is directly linked to hydrothermal activity. They are widely distributed in the Pacific Ocean and occur in the Indian, Atlantic and Arctic oceans (Gurvich, 2006; Dekov, 2008). The largest field of metalliferous sediment accumulation is in the Southeast Pacific, covering approximately 10 million km<sup>2</sup>. The second-largest fields were identified adjacent to the East Pacific Rise and at the Rodrigues Triple Junction in the Indian Ocean. Small occurrences are found in the vicinity of hydrothermal vent sites, such as Transatlantic Geotraverse, Rainbow and Grimsey, in the Mid-Atlantic Ridge. Many more are found along the mid-ocean ridges in all oceans. Accumulations are also described in back-arc settings, such as the Manus, Lau and North Fiji Basins. The significant potential of REEs plus yttrium enrichments was identified in the abyssal plain of the western North Pacific Ocean (Takaya et al., 2018). Deep-sea clays are highly enriched in REEs plus yttrium through selective adsorption processes. The Red Sea hosts 23 sites of metalliferous sediments, with Atlantis Deep II forming the largest deposit (Figure 3).



**Figure 3.** Global distribution of metalliferous sediments  
**Source:** Data compiled from Gurvich (2006) and references therein, and Dekov et al. (2008), Hrisheva & Scott (2007), Hu et al. (2017), Rusakov et al. (2013) and Pan et al. (2018) **Note:** The distribution of data records does not indicate where metalliferous sediment deposits are absent (ISA, 2021)

## RESOURCE POTENTIAL

No resource potential has been estimated for the global occurrence of metalliferous sediments. Takaya et al. (2018) estimate a rare-earth oxide resource potential of 1.2 million tons of rare-earth oxides for red REE-rich clays from the western North Pacific Ocean near Minami-Tori-Shima island. Estimates are based on a limited number of sediment cores from an area of 105 km<sup>2</sup>, down to a maximum depth of 10 m subseafloor, with extrapolated concentration maps.

The only significant resource potential has been estimated for the Atlantis II Deep basin in the Red Sea. Manafai International reassessed the resource estimate, which comprises approximately 89 million tons of dry salt-free ore, containing 3.75 million tons of zinc, 0.81 million tons of copper, 4 million tons of manganese, 5,000 tons of cobalt, 7,000 tons of silver and 47 tons of gold. The deposit forms in a complex basin adjacent to the median rift, where seafloor spreading has operated during the last 5 Ma. Exhalation of metal-rich hydrothermal brines from rift-related and cross-cutting fracture zones produced up to 25 m of metal-rich sediments over an area exceeding 57 km<sup>2</sup> (Hamer, 2017).

## TECHNOLOGY

Marine metalliferous sediments occur in different areas and geological settings on the seabed. Different conditions of formation, enrichment and occurrence may require different technological approaches for identification, exploration and potential future recovery. Metalliferous sediments associated with marine PMS sites can be identified using similar exploration strategies. Mapping and sonar devices are the primary technologies for identifying depressions that can act like sediment traps. Hydrocast profiles identify geochemical anomalies and venting in the water column and coring. Trench-cutting equipment is used to assess the thickness and subsurface composition of the sediments. Oceanographic and sedimentation processes control the distribution of metalliferous sediments on the abyssal plain and require a dense network of regional-to-local surveys, similar to those used in PMN exploration. It







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